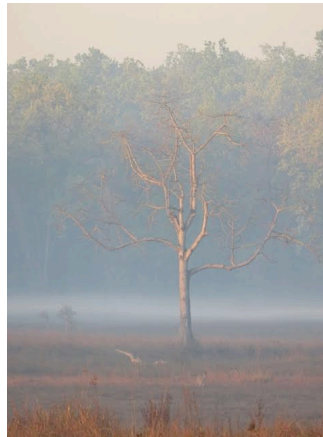


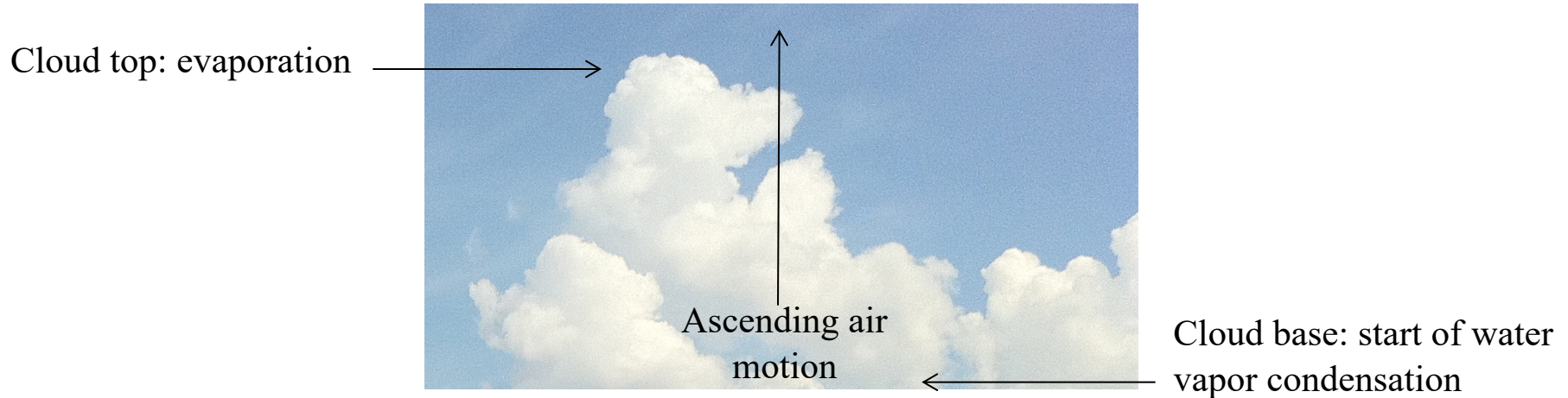
# Clouds and Acid Rain

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- General considerations on clouds and fogs
- Aqueous-phase chemistry
- Aqueous-phase chemical transformations
- Emission control strategies for acid rain



# Clouds



A cloud is formed when water vapor condenses on ultrafine particles (Aitken nuclei) and fine particles, and next on cloud droplets and ice crystals. Condensation occurs when the ascending air parcel cools adiabatically, which leads to a lower capacity to contain water vapor (i.e., increase in relative humidity): cloud base. When temperature and pressure are such that the relative humidity becomes less than 100 %, the cloud evaporates (cloud top). Therefore, there is constant air motion within a cloud.

# Cumulus Clouds

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Cumulus clouds are formed in presence of a strong vertical air motion (convection).

# Stratus Clouds

---



Stratus clouds are formed when a moist air parcel rises and, while cooling adiabatically, leads to the condensation of water vapor (since cold air has a lower capacity for water vapor content than warm air).

# Precipitating Clouds: Nimbus Cumulonimbus and Nimbostratus

---

When cloud droplets increase in size, their fall velocity increases. If the cloud droplet fall velocity becomes greater than the vertical velocity of the ascending air parcel, the cloud droplet becomes a rain drop.

- A precipitating cumulus cloud is called a cumulonimbus.
- A precipitating stratus cloud is called a nimbostratus.

# Precipitating and Non-precipitating Clouds

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Non-precipitating cumulus cloud

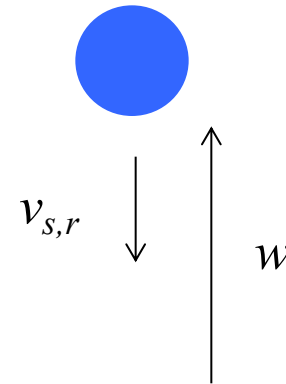


Precipitating cumulonimbus

# Precipitating and Non-precipitating Clouds



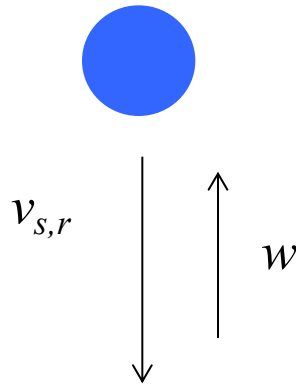
Non-precipitating cumulus cloud



The fall velocity of the cloud droplets,  $v_{s,r}$ , is less than the vertical velocity of the ascending air,  $w$ .



# Precipitating and Non-precipitating Clouds



The fall velocity of the cloud droplets (raindrops),  $v_{s,r}$ , is greater than the vertical velocity of the ascending air,  $w$ .



Precipitating cumulonimbus



# Fogs

Fog is a cloud that is in contact with the Earth's surface (visual range < 1 km)



Advection fog



Radiation fog

# Fogs

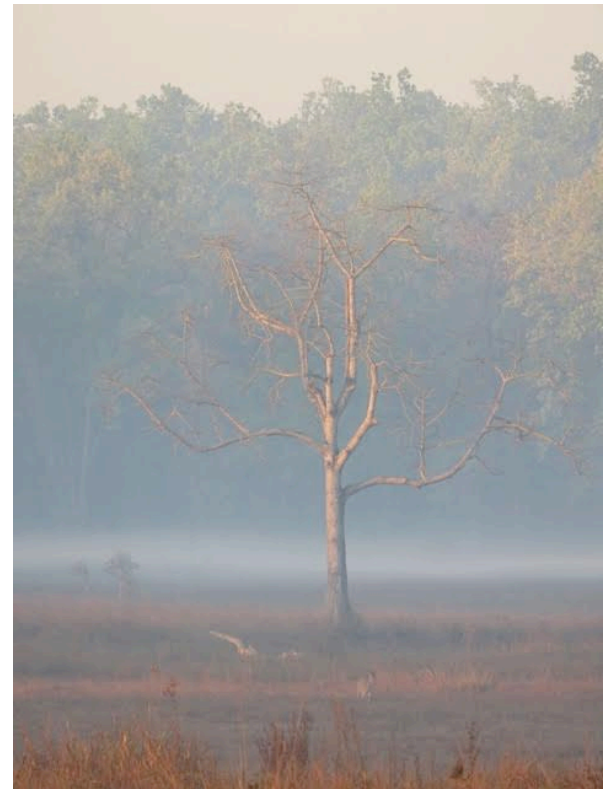


An advection fog forms when a moist air parcel gets into contact with a cold surface, which leads to a decrease in the saturation vapor pressure of water vapor and its condensation (for example, summer fog in San Francisco, California, due to the upwelling coastal cold water of the Pacific Ocean).

Advection fog

# Fogs

A radiation fog forms when the ambient temperature of a moist air parcel decreases sufficiently (for example at night), thereby leading to a decrease in the saturation vapor pressure of water vapor such that it reaches 100 % relative humidity (for example wintertime fogs in humid areas).



Radiation fog

# Clouds and Fogs

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- Liquid water contents (order of magnitude)
  - Cumulus:  $\sim 1 \text{ g/m}^3$
  - Cumulonimbus:  $> 1 \text{ g/m}^3$
  - Stratus:  $\sim 0.1 \text{ g/m}^3$
  - Nimbostratus:  $> 0.1 \text{ g/m}^3$
  - Fog:  $< 0.1 \text{ g/m}^3$
- Drops and droplets (order of magnitude of the mean diameter)
  - Non-precipitating cloud:  $\sim 40 \text{ }\mu\text{m}$
  - Rain:  $\sim 1 \text{ mm}$
  - Fog:  $\sim 10 \text{ }\mu\text{m}$

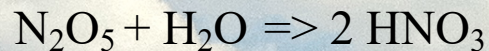
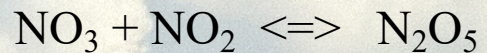
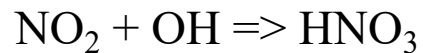
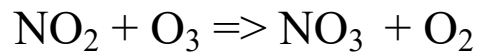
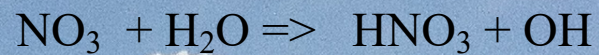
# Atmospheric Particles

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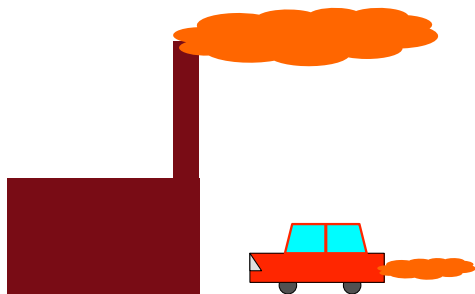
- Some particles (or fractions of particles) are formed by **chemical reaction** in the atmosphere from gases (sulfate, nitrate, ammonium, organic compounds...); these reactions can occur in the gas phase or **in the aqueous phase**: these are secondary particles or secondary PM.
- If the cloud evaporates, then some of the droplet chemical species form an atmospheric particle.

# Nitric Acid Formation

Heterogeneous reactions in clouds



$\text{NO}_x$



# Nitric Acid Formation

## Daytime Gas Phase: OH

---



These reactions are negligible at night because OH is not produced in the absence of sunlight (with the exceptions of alkene oxidation by O<sub>3</sub> and PAN reaction in presence of very high NO concentrations).

This reaction is not linear for  $\text{NO}_2 \Rightarrow \text{HNO}_3$

A change of  $x$  % in NO<sub>2</sub> leads to a change in OH radical concentrations, which affects the reaction rate and, therefore, a change in nitric acid, which is not exactly  $x$  %.

$$k = 13600 \text{ ppm}^{-1} \text{ min}^{-1}; v_r = k [\text{NO}_2] [\text{OH}]$$

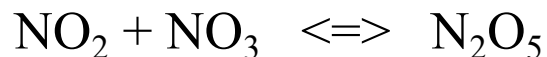
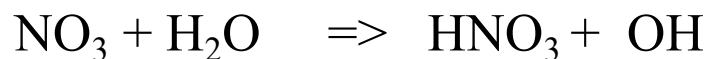
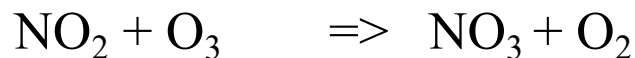
$[\text{OH}] > 10^6 \text{ cm}^{-3} \Rightarrow$  the oxidation rate of NO<sub>2</sub> is on the order of 10 % per hour during a summer day.



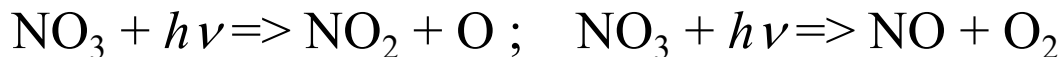
# Nitric Acid Formation

## Nighttime Gas Phase: O<sub>3</sub>

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These reactions are negligible during the day because NO<sub>3</sub> is rapidly photolyzed:

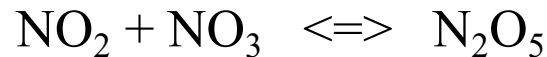
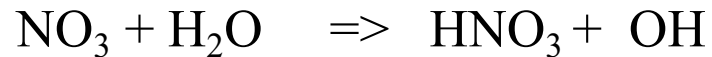
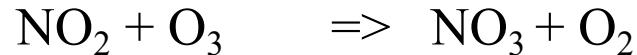


The hydrolysis reactions (NO<sub>3</sub> + H<sub>2</sub>O and N<sub>2</sub>O<sub>5</sub> + H<sub>2</sub>O) are much faster when they occur heterogeneously at the surface of droplets.

# Nitric Acid Formation

## Nighttime Gas Phase: O<sub>3</sub>

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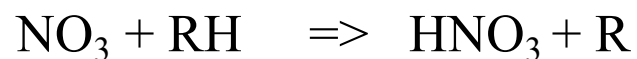
This system of reactions is strongly non-linear for  $\text{NO}_2 \Rightarrow \text{HNO}_3$

A change of  $x$  % in  $\text{NO}_2$  leads to a change in  $\text{O}_3$  concentrations, which affects the rate of the first reaction and, therefore, leads to a change in nitric acid, which is not exactly  $x$  % (it can be greater or less than  $x$ ).

# Nitric Acid Formation

## Gas phase: Organic Compounds

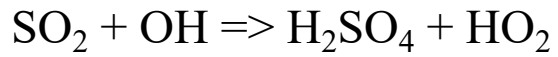
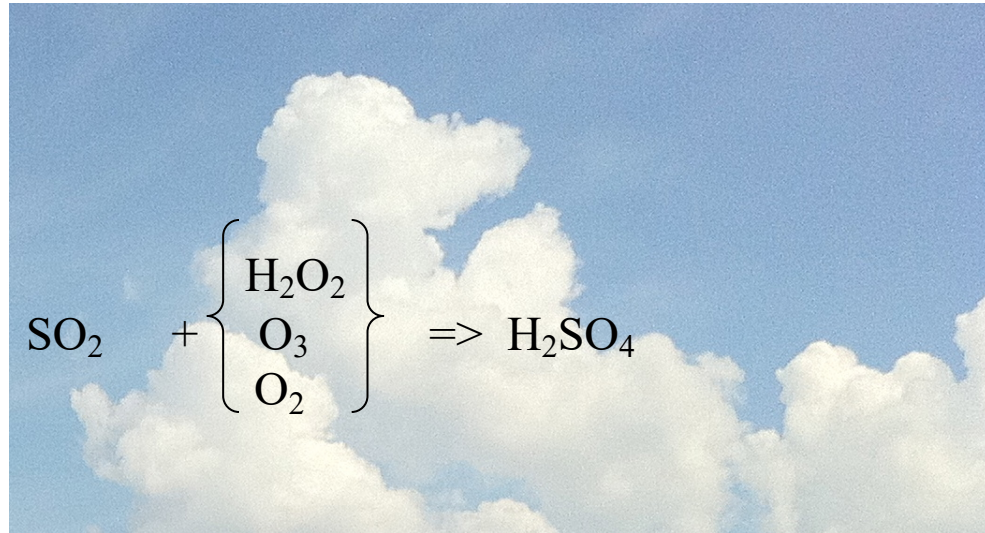
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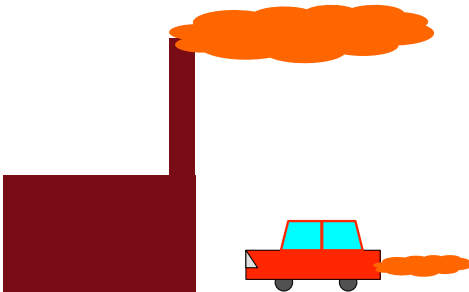
These reactions are generally minor for the formation of HNO<sub>3</sub>.

# Sulfuric Acid Formation

System of non-linear reactions in clouds



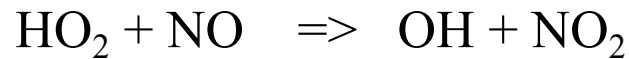
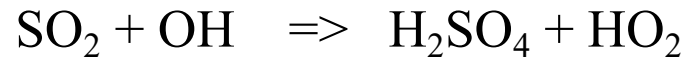
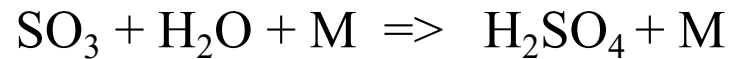
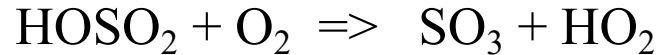
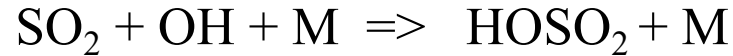
SO<sub>2</sub>



# Sulfuric Acid Formation

## Gas Phase: OH

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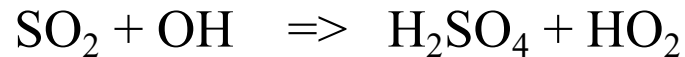
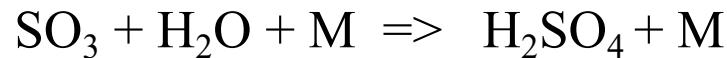
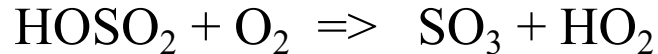
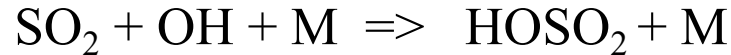


This system of reaction is linear for  $\text{SO}_2 \Rightarrow \text{H}_2\text{SO}_4$

A change of  $x$  % in  $\text{SO}_2$  leads to a change of  $x$  % in sulfuric acid.

# Sulfuric Acid Formation

## Gas Phase: OH



The first reaction governs the overall reaction rate because the following reactions are much faster. Therefore, the overall kinetics is governed by the rate of the first reaction:

$$k = 1400 \text{ ppm}^{-1} \text{ min}^{-1} ; v_r = k [\text{SO}_2] [\text{OH}]$$

$[\text{OH}] > 10^6 \text{ cm}^{-3} \Rightarrow$  the oxidation rate of  $\text{SO}_2$  is on the order of 1 % per hour during a summer day.

# Atmospheric Aqueous Phase Concentration and Activity

---

Concentration of a chemical species in the aqueous phase:

- Molar fraction: number of moles per mole of water ( $x_i$ )
- Molality: number of moles per kg of water ( $m_i$ )
- Molarity: number of moles per liter of water ( $c_i$ )

Therefore, for trace pollutants:  $m_i \sim c_i$

Activity of a chemical species in a non-ideal solution:

$$\alpha_i = \gamma_i x_i$$

where:

$x_i$  = molar fraction

$\alpha_i$  = activity

$\gamma_i$  = activity coefficient; it represents the deviation of the solution with respect to an ideal solution.



# Atmospheric Aqueous Phase Concentration and Activity

---

Activity of a chemical species in a non-ideal solution:

$$\alpha_i = \gamma_i x_i$$

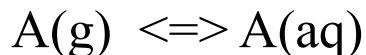
For a cloud droplet, one may assume that the solution is ideal because the concentrations of chemical species (other than water) are small:

$$\gamma_i = 1 \text{ and } \alpha_i = x_i$$

However, for a small particle, the solution will not be ideal because the concentrations of chemical species will be high.

# Atmospheric Aqueous Phase Equilibrium with the Gas Phase

---



Henry's law (dilute solution):

$$[A(aq)] = H_A P_A$$

$[A(aq)]$ : molarity of A (moles per liter or M)

$P_A$ : partial pressure of A (atm)

$H_A$ : Henry's law constant (M atm<sup>-1</sup>)

Important points:

- The Henry's law constant can be defined with other units and is sometimes defined as the inverse ratio (concentration in the gas phase / concentration in the aqueous phase).
- Henry's law is valid only for species present in low concentrations.

# Effective Henry's Law Constant

## Monoacid: HNO<sub>3</sub>

---

Gas/water equilibrium     HNO<sub>3</sub>(g)  $\rightleftharpoons$  HNO<sub>3</sub>(aq)

Dissociation equilibrium     HNO<sub>3</sub>(aq)  $\rightleftharpoons$  NO<sub>3</sub><sup>-</sup> + H<sup>+</sup>

Equilibrium constants:

$$[\text{HNO}_3(\text{aq})] = H_{\text{HNO}_3} P_{\text{HNO}_3}$$
$$[\text{NO}_3^-] [\text{H}^+] / [\text{HNO}_3(\text{aq})] = K_{\text{HNO}_3}$$

[HNO<sub>3</sub>(aq)], [NO<sub>3</sub><sup>-</sup>], [H<sup>+</sup>]: molarities (M)

$P_{\text{HNO}_3}$ : partial pressure of NO<sub>3</sub> (atm)

$H_{\text{HNO}_3}$ : Henry's law constant (M atm<sup>-1</sup>)

$K_{\text{HNO}_3}$ : dissociation constant of acid HNO<sub>3</sub> (M)

# Effective Henry's Law Constant

## Monoacid: HNO<sub>3</sub>

---

One defines the effective Henry's law constant,  $H_{\text{HNO}_3, \text{eff}}$ , as the equilibrium between the species in the gas phase and all corresponding species in the aqueous phase:

$$[\text{HNO}_3(\text{aq})] + [\text{NO}_3^-] = H_{\text{HNO}_3, \text{eff}} P_{\text{HNO}_3}$$

Substituting:

$$[\text{HNO}_3(\text{aq})] + ([\text{HNO}_3(\text{aq}) K_{\text{HNO}_3} / [\text{H}^+]) = H_{\text{HNO}_3, \text{eff}} P_{\text{HNO}_3}$$

$$[\text{HNO}_3(\text{aq})] (1 + K_{\text{HNO}_3} / [\text{H}^+]) = H_{\text{HNO}_3, \text{eff}} P_{\text{HNO}_3}$$

$$H_{\text{HNO}_3, \text{eff}} = H_{\text{HNO}_3} (1 + K_{\text{HNO}_3} / [\text{H}^+])$$

# Effective Henry's Law Constant

## Diacid: Sulfuric Acid (H<sub>2</sub>SO<sub>4</sub>)

---

This equation can be generalized in the case of a second dissociation (for diacids such as sulfuric acid).



Dissociation equilibria



Effective Henry's law constant:

$$H_{\text{H}_2\text{SO}_4, \text{eff}} = H_{\text{H}_2\text{SO}_4} (1 + (K_1 / [\text{H}^+]) + (K_1 K_2 / ([\text{H}^+]^2)))$$

# Effective Henry's Law Constants Bases and Acids

---

One can also define effective Henry's law constants for bases (for example, for ammonia,  $\text{NH}_3$ ).

The effective Henry's law constant depends on pH.

For an acid, it decreases as the pH decreases (less dissociation of acids in an acidic solution).

For a base, it increases as the pH decreases (more dissociation of bases in an acidic solution).

# Solubility of some Air Pollutants

Chemical species	Henry's law constant (M atm <sup>-1</sup> )	Effective Henry's law constant <sup>a</sup> (M atm <sup>-1</sup> )	Fraction in the aqueous phase
NO	$1.9 \times 10^{-3}$	$1.9 \times 10^{-3}$	0.000005 %
NO <sub>2</sub>	$1.2 \times 10^{-2}$	$1.2 \times 10^{-2}$	0.00002 %
O <sub>3</sub>	$1.1 \times 10^{-2}$	$1.1 \times 10^{-2}$	0.00002 %
CO <sub>2</sub>	$3.4 \times 10^{-2}$	$4 \times 10^{-2}$	0.0001 %
SO <sub>2</sub>	1.23	$6.5 \times 10^3$	14 %
NH <sub>3</sub>	62	$2.6 \times 10^5$	87 %
HCHO <sup>b</sup>	$6.3 \times 10^3$	$6.3 \times 10^3$	13 %
H <sub>2</sub> O <sub>2</sub>	$7.45 \times 10^4$	$7.45 \times 10^4$	65 %
HNO <sub>3</sub>	$2.1 \times 10^5$	$1.3 \times 10^{12}$	100 %

$H < 10^3 \text{ M atm}^{-1}$   
not soluble

$10^3 < H < 10^5$   
soluble

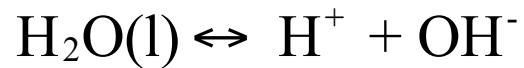
$H > 10^5$   
very soluble

(a) pH = 5.6; LWC = 1 g/m<sup>3</sup>; T = 25 ° C; (b) including diol formation



# Dissociation Equilibrium of Water

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$\text{H}^+$  is the hydrogen ion (proton);

$\text{OH}^-$  is the hydroxide ion

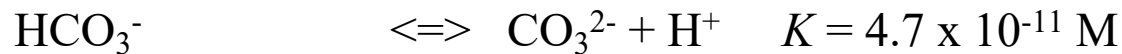
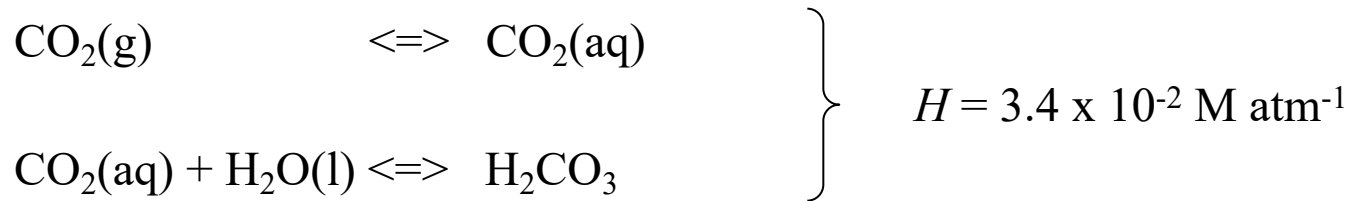
$$K'_{\text{H}_2\text{O}} = \frac{[\text{H}^+][\text{OH}^-]}{[\text{H}_2\text{O}(\text{l})]} = 1.82 \times 10^{-16} \text{ M}$$

$$K_{\text{H}_2\text{O}} = [\text{H}^+][\text{OH}^-] = 10^{-14} \text{ M}^2$$

$$\text{pH} = -\log(\gamma_{\text{H}^+} [\text{H}^+])$$

pH = 7.0 for pure water at 298 K

# Atmospheric Aqueous Phase



Electroneutrality equation of the aqueous phase:

$$[\text{H}^+] = [\text{HCO}_3^-] + 2 [\text{CO}_3^{2-}] + [\text{OH}^-]$$

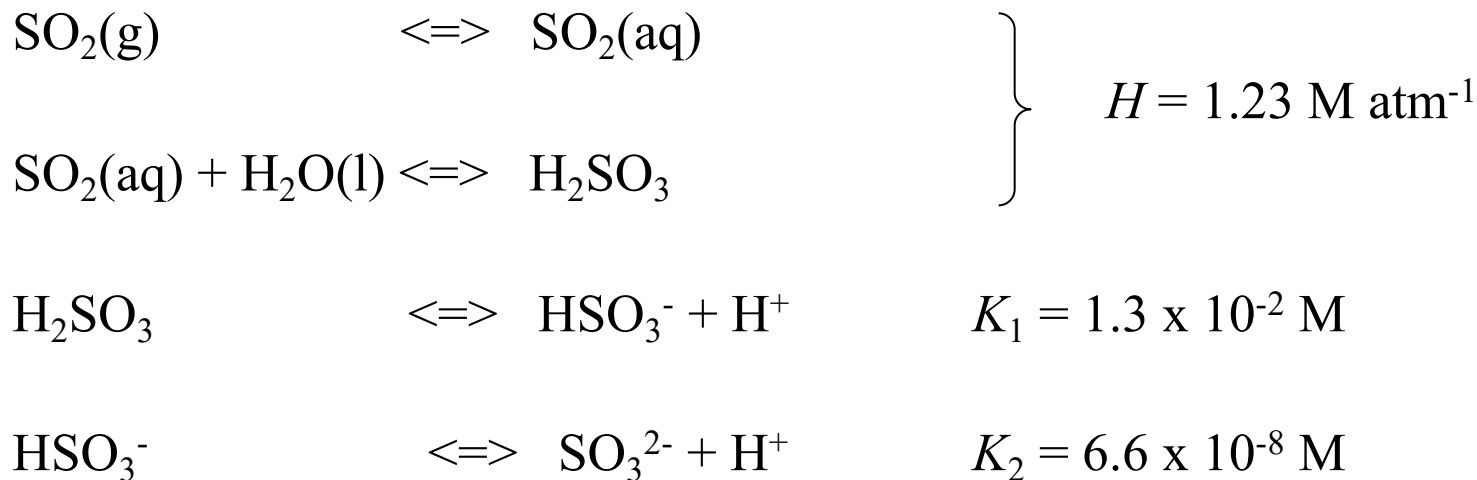
# Atmospheric Aqueous Phase

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The concentration of carbon dioxide is currently about 400 ppm

The pH of water in clouds can be calculated from the atmospheric concentration of CO<sub>2</sub>: it is about 5.6.

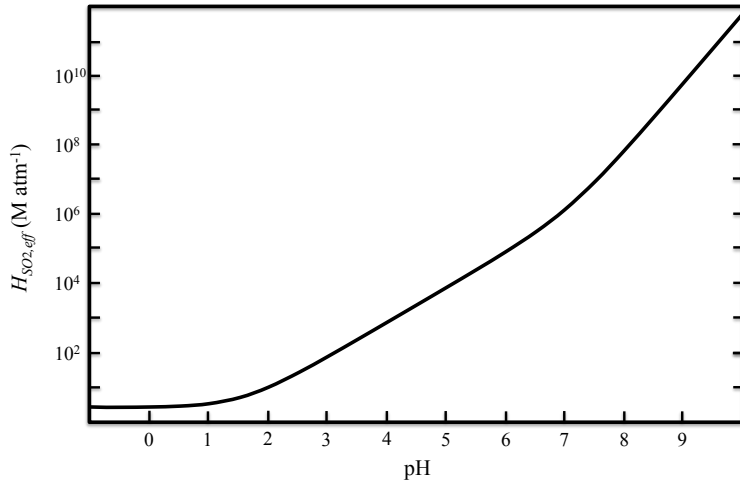
# Dissolution of SO<sub>2</sub> in the Aqueous Phase



SO<sub>2</sub> is a weak acid: it is not completely dissociated in the aqueous phase

$$[\text{S(IV)}] = [\text{H}_2\text{SO}_3] + [\text{HSO}_3^-] + [\text{SO}_3^{2-}]$$

# Dissolution of SO<sub>2</sub> in the Aqueous Phase



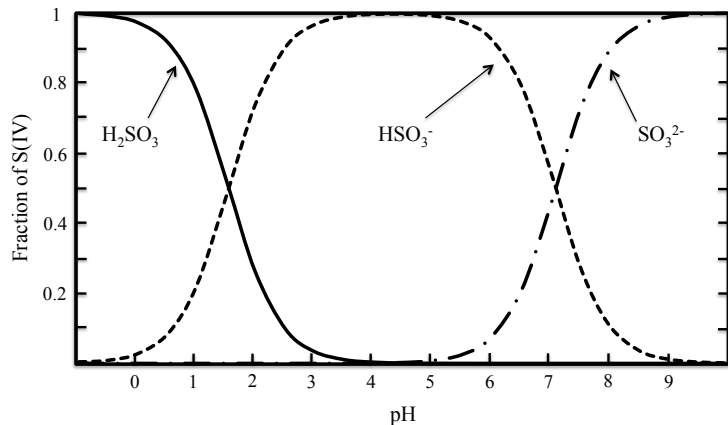
$$[S(IV)(aq)] = H_{eff} P_{SO_2}$$

$$H_{eff} = H \left( 1 + \frac{K_1}{[H^+]} + \frac{K_1 K_2}{[H^+]^2} \right)$$

$$\frac{[H_2SO_3(aq)]}{[S(IV)(aq)]} = \left( 1 + \frac{K_1}{[H^+]} + \frac{K_1 K_2}{[H^+]^2} \right)^{-1}$$

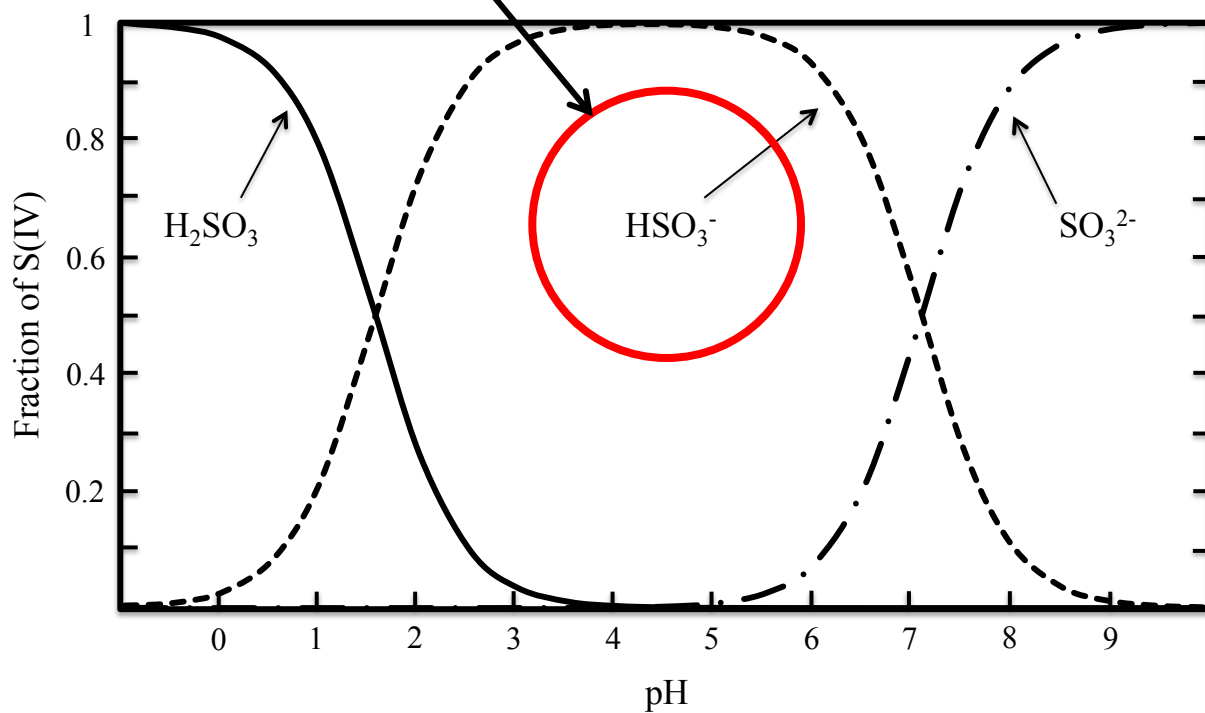
$$\frac{[HSO_3^-(aq)]}{[S(IV)(aq)]} = \left( 1 + \frac{[H^+]}{K_1} + \frac{K_2}{[H^+]} \right)^{-1}$$

$$\frac{[SO_3^{2-}(aq)]}{[S(IV)(aq)]} = \left( 1 + \frac{[H^+]^2}{K_1 K_2} + \frac{[H^+]}{K_2} \right)^{-1}$$



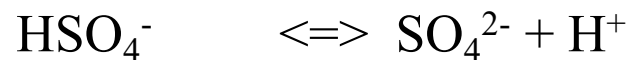
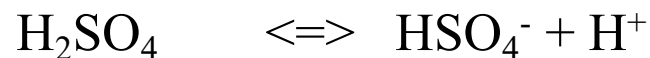
# Dissolution of SO<sub>2</sub> in the Aqueous Phase

Range of pH values in atmospheric clouds: HSO<sub>3</sub><sup>-</sup> is the major species



# Dissolution of H<sub>2</sub>SO<sub>4</sub> in the Aqueous Phase

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H<sub>2</sub>SO<sub>4</sub> is a strong acid: it dissociates easily in the aqueous phase.

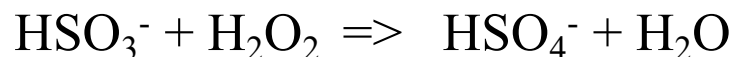
$$[\text{S(VI)}] = [\text{H}_2\text{SO}_4] + [\text{HSO}_4^-] + [\text{SO}_4^{2-}]$$



# Sulfuric Acid Formation

## Aqueous Phase: H<sub>2</sub>O<sub>2</sub>

---



This reaction does not depend much on pH.

The rate of this reaction is fast and the lifetime of SO<sub>2</sub> is on the order of only a few minutes: it is a titration reaction.

$$H = 7.4 \times 10^4 \text{ M / atm for H}_2\text{O}_2$$

$$\frac{d[\text{H}_2\text{SO}_4]}{dt} = 7.2 \times 10^7 [\text{H}^+] [\text{S(IV)}] [\text{H}_2\text{O}_2]$$

M s<sup>-1</sup>

# Sulfuric Acid Formation

## Aqueous Phase: H<sub>2</sub>O<sub>2</sub>

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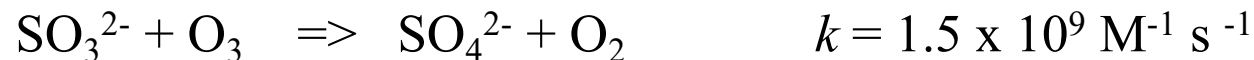
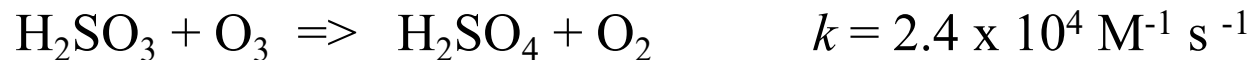


If  $[\text{SO}_2] < [\text{H}_2\text{O}_2]$ : a change of  $x$  % in  $\text{SO}_2$  leads to a change of  $x$  % in sulfuric acid and the system is linear.

If  $[\text{SO}_2] > [\text{H}_2\text{O}_2]$ : a change of  $x$  % in  $\text{SO}_2$  leads to a change of 0 % in sulfuric acid as long as  $[\text{SO}_2] > [\text{H}_2\text{O}_2]$  and the system is non-linear; the reaction stops when  $\text{H}_2\text{O}_2$  has been totally consumed: one says that the system is oxidant-limited.

# Sulfuric Acid Formation

## Aqueous Phase: O<sub>3</sub>



$$H_{\text{O}_3} = 1.1 \times 10^{-2} \text{ M/atm}$$

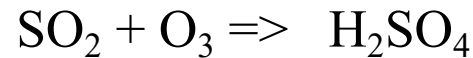
This reaction depends strongly on pH.

The overall reaction rate is fast but the lifetime of SO<sub>2</sub> is on the order of several tens of minutes; therefore, it is typically slower than the oxidation by H<sub>2</sub>O<sub>2</sub>.

# Sulfuric Acid Formation

## Aqueous Phase: O<sub>3</sub>

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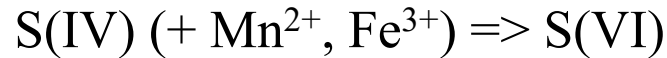
When H<sub>2</sub>SO<sub>4</sub> is formed, the pH decreases (more acidic solution), the rate slows down and H<sub>2</sub>SO<sub>4</sub> formation decreases => this reaction is self-limiting.

Therefore, a decrease of  $x$  % in SO<sub>2</sub> does not lead to a change of  $x$  % in sulfuric acid, but to a lesser decrease.

# Sulfuric Acid Formation

## Aqueous Phase: O<sub>2</sub>

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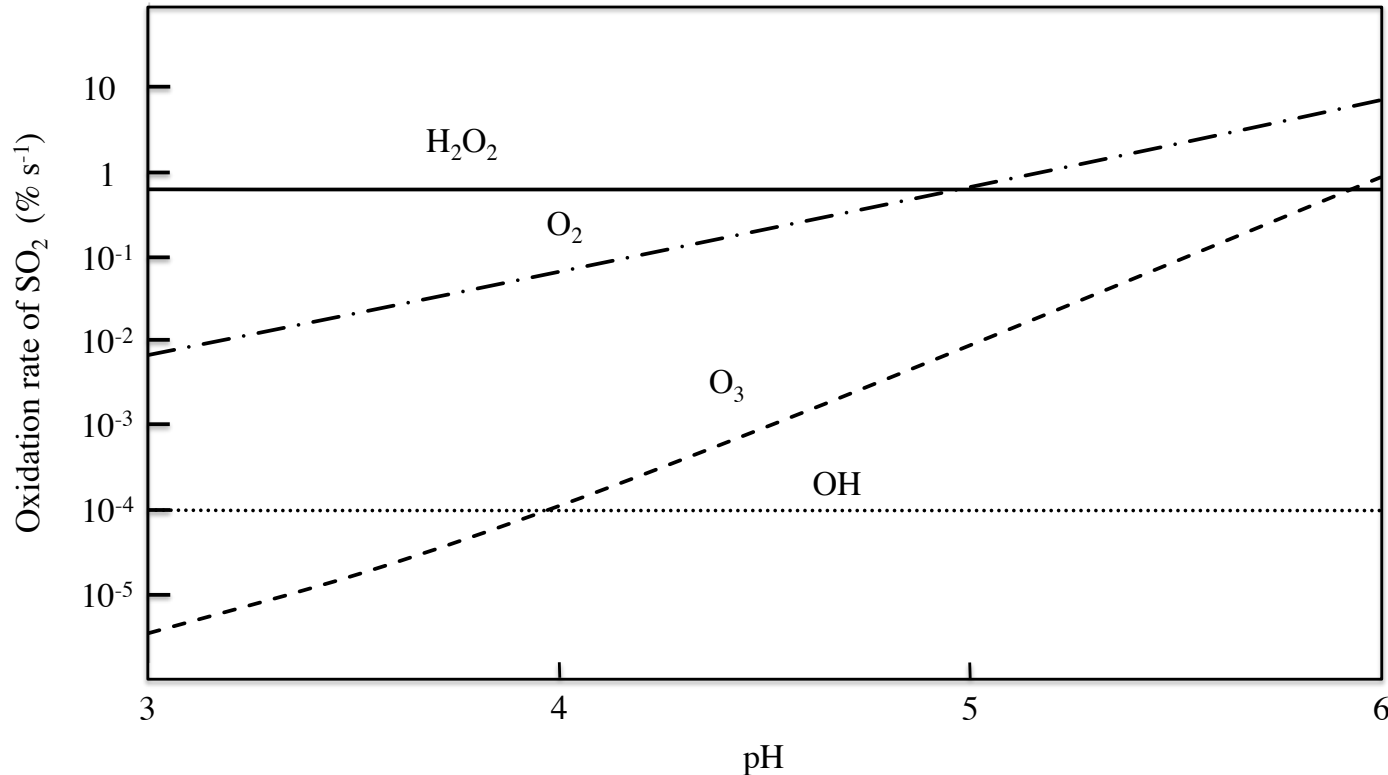
This reaction is catalyzed by metal ions such as Fe<sup>3+</sup> and Mn<sup>2+</sup>.

The overall reaction rate is fast only when pH > 4; the kinetics slows down when the pH decreases and, therefore, this reaction is self-limiting.

Therefore, a decrease of  $x$  % in SO<sub>2</sub> does not lead to a change of  $x$  % in sulfuric acid, but to a lesser decrease.

# Aqueous-phase Chemistry

## Kinetics of SO<sub>2</sub> Oxidation



Conditions for the cloud aqueous phase are as follows: the cloud is at about 3 km altitude,  $P = 0.69$  atm,  $T = 5$  ° C,  $L = 0.5$  g m<sup>-3</sup>. Conditions for the gas phase are as follows:  $P = 1$  atm,  $T = 25$  ° C, clear sky. Concentrations are as follows: [SO<sub>2</sub>(g)] = 2 ppb, [H<sub>2</sub>O<sub>2</sub>(g)] = 2 ppb, [O<sub>3</sub>(g)] = 40 ppb, [Fe(aq)] = 10 mM, [Mn(aq)] = 0,5 mM, [OH(g)] = 10<sup>6</sup> cm<sup>-3</sup>.

# Aqueous-phase Chemistry

## Kinetics of SO<sub>2</sub> Oxidation

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- In an acidic solution, oxidation by H<sub>2</sub>O<sub>2</sub> is the only important oxidation pathway.
- Oxidation by ozone and oxygen is self-limiting because the formation of sulfuric acid leads to a decrease in pH and, therefore, the kinetics slows down.
- Gas-phase oxidation by OH is very slow.

# Emission Control Strategies for Acid Rain

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- Acid rain is due mostly to sulfuric acid ( $\text{H}_2\text{SO}_4$ ) and nitric acid ( $\text{HNO}_3$ ).
- The formation of  $\text{H}_2\text{SO}_4$  is dominated by aqueous-phase chemistry because its gas-phase chemistry is slow. The reduction of  $\text{SO}_2$ , precursor of  $\text{H}_2\text{SO}_4$ , may be strongly non-linear if the  $\text{SO}_2$  concentrations are greater than the  $\text{H}_2\text{O}_2$  concentrations (i.e., in an oxidant-limited regime). It will be almost linear once the  $\text{SO}_2$  concentrations become less than the  $\text{H}_2\text{O}_2$  concentrations (some non-linearity remains due to the pH dependency of the  $\text{O}_3$  and  $\text{O}_2$  reactions).
- The reduction of the precursors of  $\text{HNO}_3$  may lead to a non-linear response of nitric acid formation, because it also affects oxidant concentrations (via the  $\text{VOC}/\text{NO}_x$  ratio of photochemical smog formation).